# Components of forest soil CO<sub>2</sub> efflux as estimated from Δ<sup>14</sup>C values of SOM

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#### Introduction

Partitioning of soil CO $_2$  efflux into autotrophic (SOM-derived) and heterotrophic (root-derived) respiration is a research priority to improve our ability to predict the net response of soil carbon stores to climate change. Using the  $\Delta^{14}$ C value of a SOM pool measured at a point in time, a turnover time of the SOM pool can be estimated using a simple model in conjunction with a record of atmospheric  $\Delta^{14}$ CO $_2$  concentrations. Following the conversion of  $\Delta^{14}$ C values to turnover times, soil CO $_2$  efflux from SOM pools can be calculated as the ratio of their carbon stocks and turnover times.

#### **Methods**

#### Site selection and soil respiration measurements

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The study was performed in northern Italy in eleven forest sites located along an elevation gradient from 220 to 1740 meters above sea level (a.s.l.). Seven sites were dominated by an evergreen tree species and the remaining four by a deciduous species (see Table 1). Soil  $\mathrm{CO}_2$  efflux was investigated in a previous study by Rodeghiero and Cescatti (2005), who determined total annual soil  $\mathrm{CO}_2$  efflux at each site using chamber based soil respiration measurements.

## **SOM Turnover Times:**

Turnover times for litter, light and heavy SOM fractions were calculated, based on their  $\Delta^{14}C$  values, by applying a modeling approach (Hakkenberg et al., 2008), which assumes that (1) annual carbon inputs to the SOM pool are constant and (2) SOM pools are at steady state. The change in  $^{14}C/^{12}C$  isotope ratio of a SOM pool (RSOM) over time can then be described as (following Gaudinski et al., 2000; Bird & Torn, 2006): RSOM (t) = k x Ratm (t) + RSOM (t - 1) x (1 - k - \lambda) where k = SOM decomposition rate (yr-¹);  $\lambda$ = $^{14}C$  radioactive decay constant (0.00012 yr-¹); Ratm =  $^{14}C/^{12}C$  isotope ratios of atmospheric CO $_2$  calculated from atmospheric  $\Delta^{14}C$  values; t = time (yr) for which calculation is being performed.

Table 1 - Some characteristics of the investigated sites

Site	Elevation	(A)Y	Dominant	Litterfall
(abbreviation)	(m a.s.l.)	Forest type	species	(gC/m2/yr)
Casteller (CaH)	220	Hop hornbeam	Ostrya carpinifolia	214 (26)
Toblino (ToO)	250	Holm oak	Quercus ilex	515 (72)
Mattarello (MaP)	420	Austrian pine	Pinus nigra	328 (24)
Lagolo P (LaP)	760	Scots pine	Pinus sylvestris	188 (6)
Lagolo B (LaB)	970	European beech	Fagus sylvatica	233 (25)
Brigolina B (BrB)	1020	European beech	Fagus sylvatica	257 (31)
Brigolina S (BrS)	1020	Norway spruce	Picea abies	275 (21)
Vaneze (VaS)	1150	Norway spruce	Picea abies	147 (17)
Lavarone (LaF)	1370	Silver fir	Abies alba	254 (12)
Bondone (BoB)	1470	European beech	Fagus sylvatica	177 (6)
Renon (ReS)	1740	Norway spruce	Picea abies	115 (12)

## References

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## Results

The total  $\Delta^{14}$ C-based decomposition fluxes of the sites ranged from 164 to 447 g C/m²/yr, with by far the largest contribution (on average 81%) from the litter layers (Figure 1).

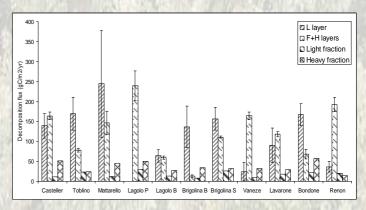


Fig. 1 – Estimated annual carbon fluxes released by the decomposition of the three SOM fractions.

The fraction of total soil  $CO_2$  efflux remaining after subtracting the total SOM-derived  $CO_2$  flux from the chamber measured flux was presumed to be respiration derived from recently fixed carbon (Figure 2). The root respiration from recently fixed carbon total  $CO_2$  efflux and the estimated respiration of the litter within the first year after litterfall . This resulted in an average root-derived respiration of  $\sim 50\%$  of the total soil  $CO_2$  efflux.

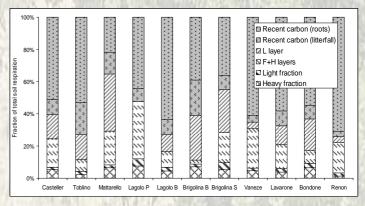


Fig. 2 – Partitioning of total annual soil CO<sub>2</sub> efflux into SOM-derived components and components derived from recently fixed carbon.

## Conclusions

The decomposition flux derived from the litter layers was by far the largest fraction of the total decomposition flux (on average 81%).

The Δ¹4C-based decomposition fluxes of the different SOM fractions as estimated here, suggested that most of the annual carbon inputs to the SOM pool took place aboveground and to a much smaller degree in the mineral soil.

A very large fraction of the total soil  ${\rm CO_2}$  efflux (~ 90%) resulted from root respiration and decomposition of fresh organic matter ( $\Delta^{14}{\rm C}$ -based litter-derived respiration plus the respiration from recently fixed carbon).

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